

# Faster Tropical Upper Stratospheric Upwelling Drives Changes in Ozone Chemistry

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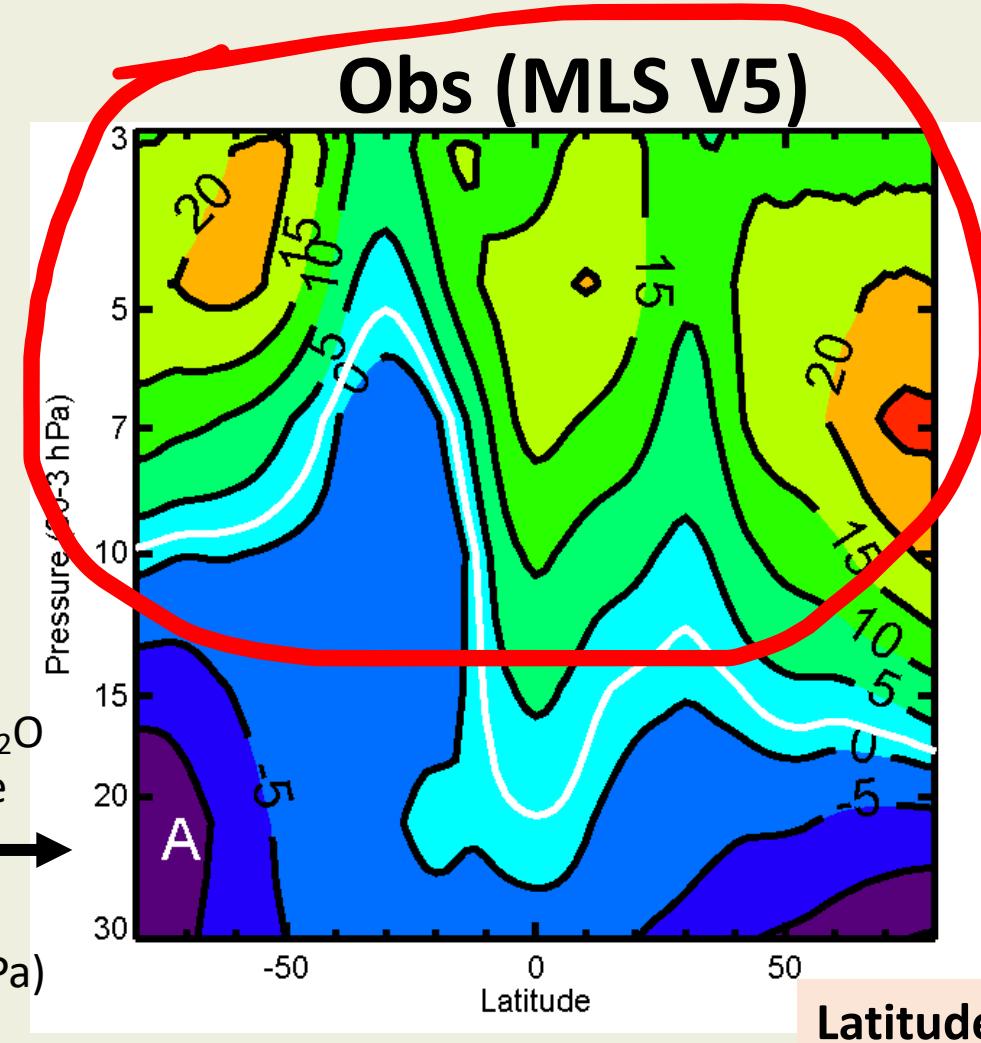
**BOTTOM LINE: Upper Strat Composition Change →  
*O<sub>3</sub> Loss Changes → O<sub>3</sub> trends from 2005-2021***

# Quasi-Decadal (QD) N<sub>2</sub>O Change over 2005-2021 *(Percent change of 2013-2021 mean - 2005-2013 mean)*

N<sub>2</sub>O surface  
growth over this  
period = 2%

No drift in MLS N<sub>2</sub>O  
22 hPa and above

Known negative  
drift below 22 hPa

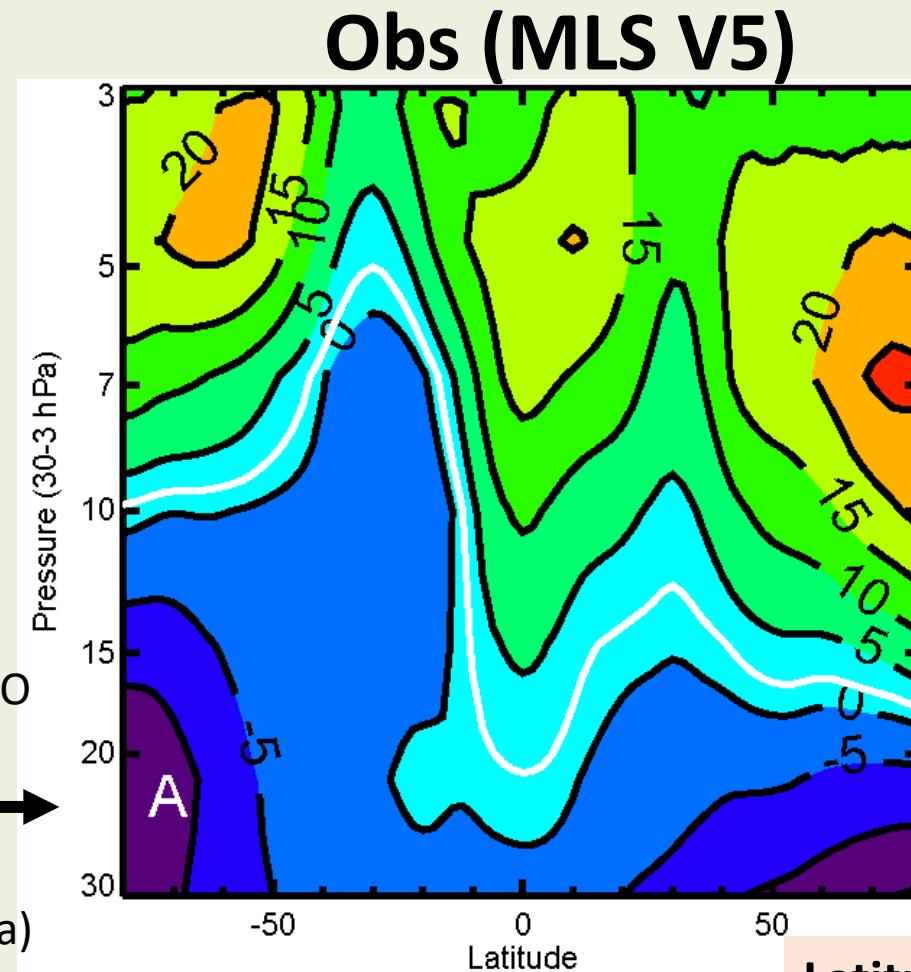


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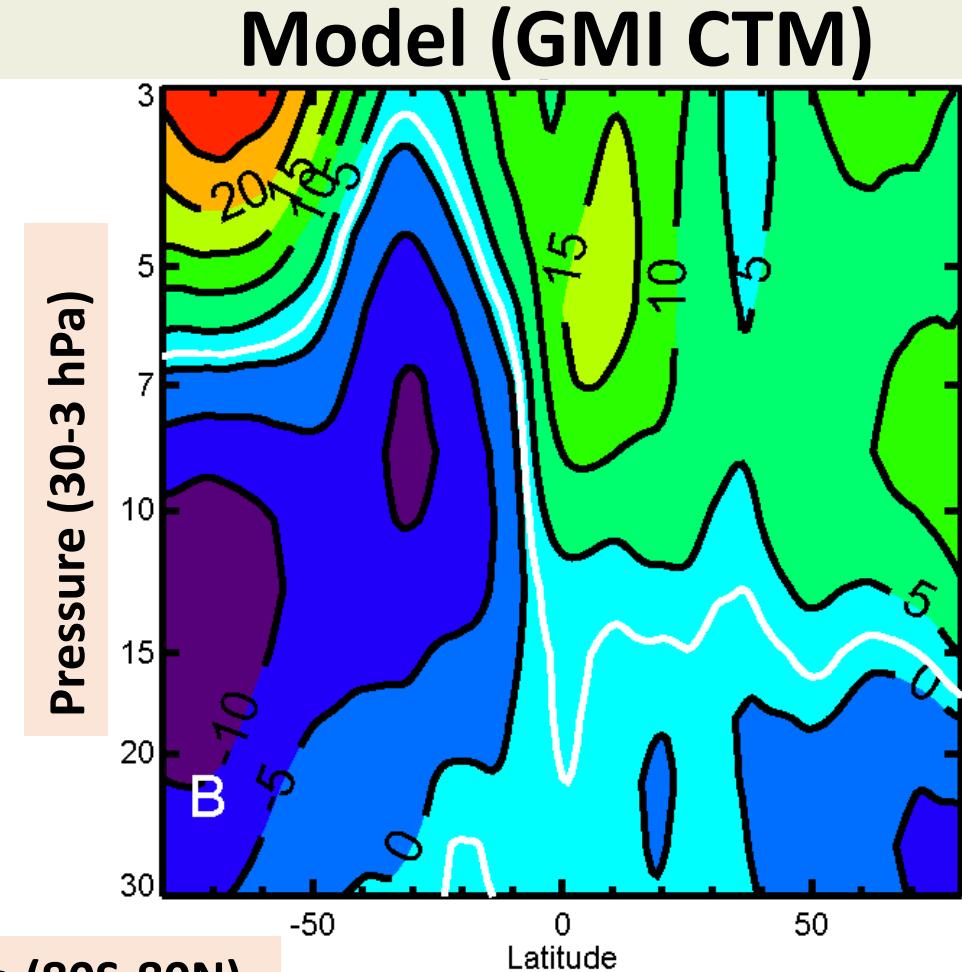
N<sub>2</sub>O surface growth over this period = 2%  
(shown by white contour)

No drift in MLS N<sub>2</sub>O  
22 hPa and above

Known negative drift below 22 hPa

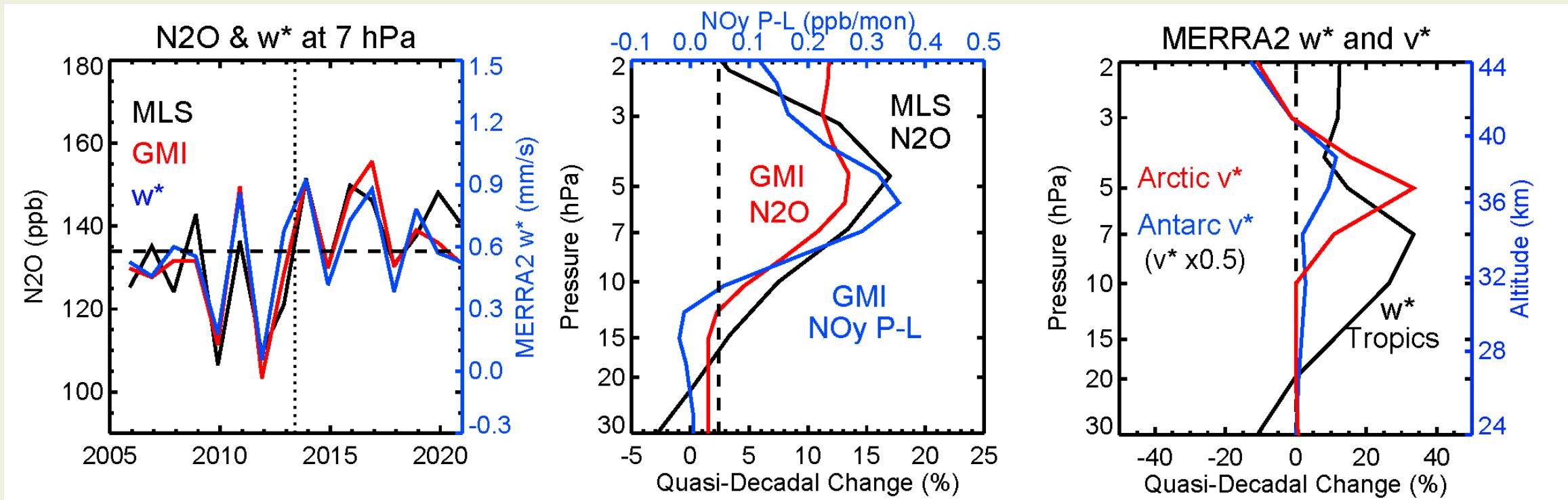


Latitude (80S-80N)



# Upper Stratospheric Circulation Change ( $w^*$ and $v^*$ )

Explains why is  $N_2O$  so high



Tropical  $N_2O$  grows above 10 hPa, leading to increased odd nitrogen production.

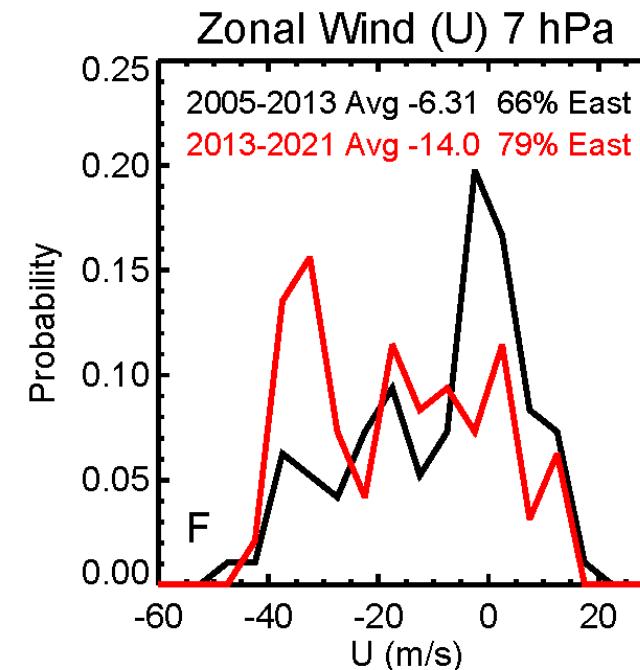
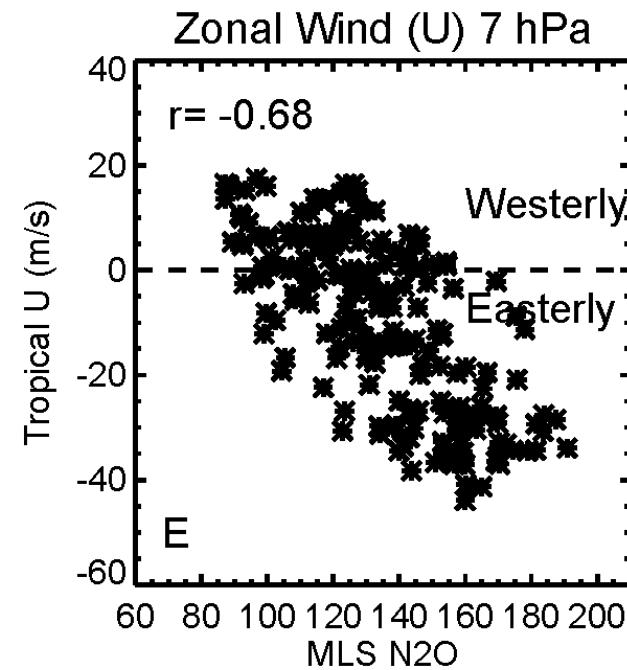
Upwelling ( $w^*$ ) increases above 20 hPa up to 30%  
Arctic poleward transport ( $v^*$  50-70°N) increases by 60% at 5 hPa

# The Quasi-Biennial Oscillation (QBO) modulates tropical upwelling: stronger during Easterlies, weaker in Westerlies

MLS tropical monthly mean  $\text{N}_2\text{O}$  is positively correlated with  $w^*$  and anticorrelated with the zonal wind (i.e., QBO)

Stronger and more frequent QBO Easterlies 10-3 hPa increased mean tropical upwelling from 2013-2021

MLS and MERRA2 Monthly data from June 2005 to May 2021



2013-2021 was QBO-E  
79% of the time

# Two GMI Chemistry Transport Model (CTM) simulations with the same chemistry but different dynamics

GMI Chemistry Transport Model integrated with MERRA2:

- BASELINE: Time-varying MERRA2 Fields from June 2005-May 2021
- Fixed Dyn: June 2005-May 2007 Merra2 Fields (2 years) recycled until May 2021
- Both Simulations are forced with the same time-dependent source gases ( $\text{N}_2\text{O}$ , CFCs)

These start/end dates for the 2-year repeat chosen because of a 2-yr period QBO. This minimizes the transport adjustment when recycling.

$\text{O}_3$  loss differences between BASELINE and FIXED DYN show how *dynamical changes led to chemical changes* that affected  $\text{O}_3$ .

# Compare QD Change in $O_3$ Loss from $NO_x$ and $ClO_x$ (ppb $O_3/mo$ ) in BASELINE and Fixed Dyn Simulations

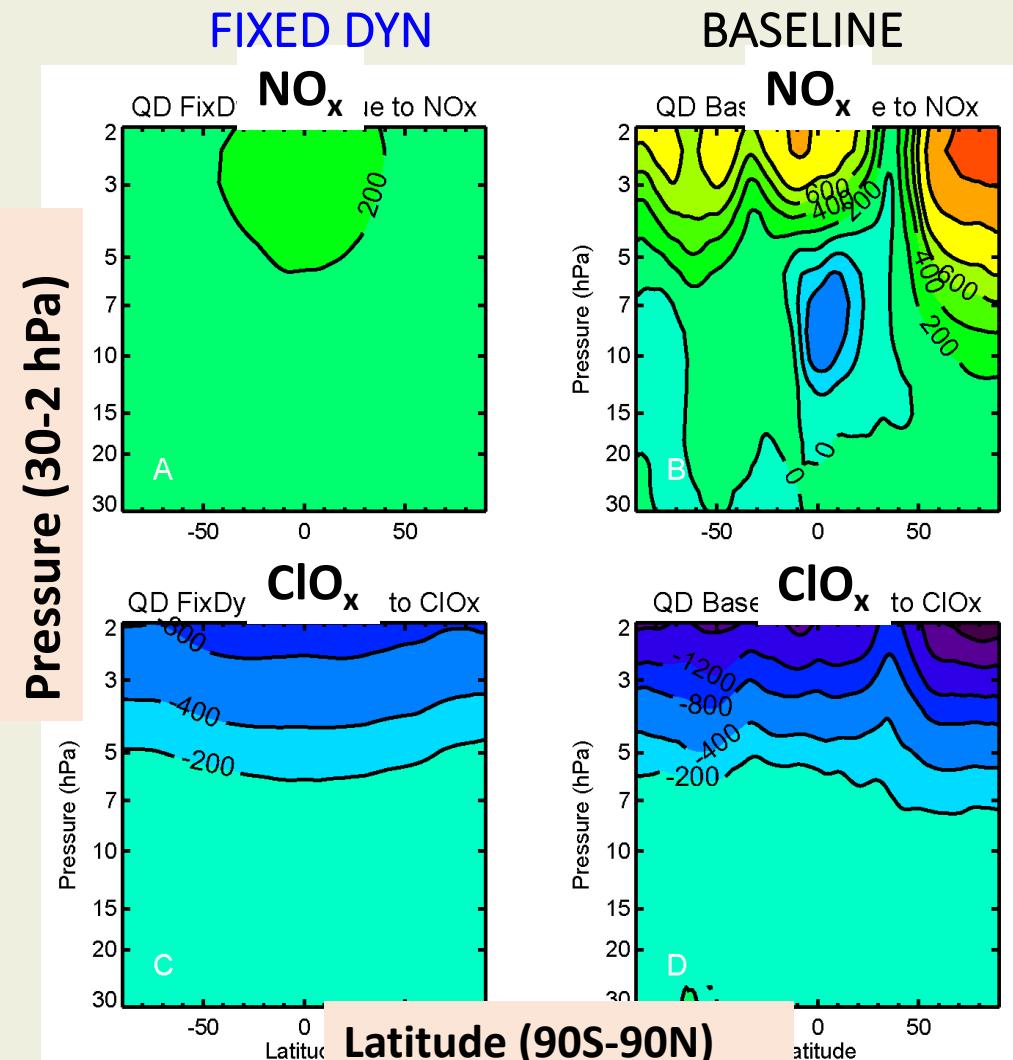
Fixed Dyn:

Hemispheric Symmetry

This is consistent with expectations from trends in tropospheric source gases (CFCs,  $N_2O$ )

Loss by  $NO_x$  increases above 5 hPa

Loss by  $ClO_x$  decreases with increasing altitude



BASELINE:

*Hemispheric asymmetry*

Greater loss by  $NO_x$  above 5 hPa, especially in the Arctic

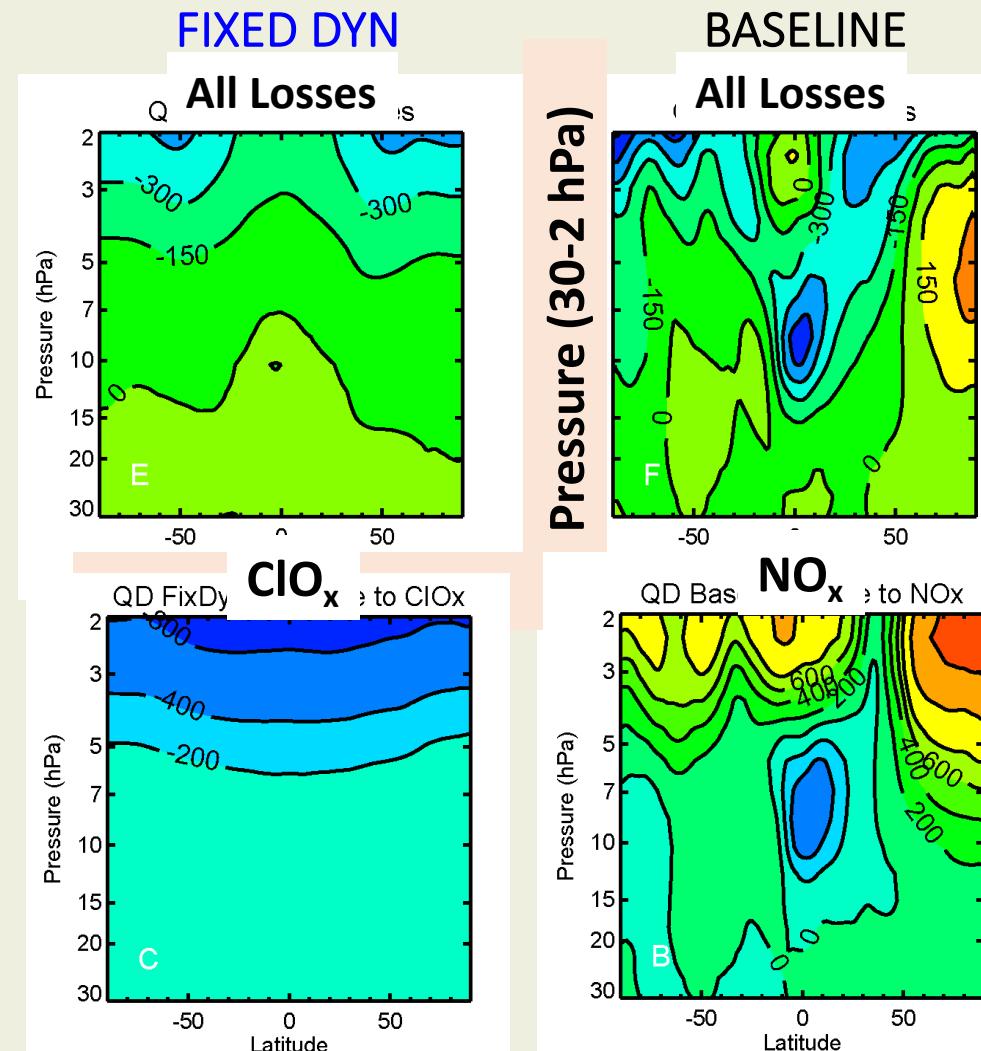
Less loss by  $NO_x$  in the tropical middle stratosphere

Loss by  $ClO_x$  decreases even more than FixedDyn in the Arctic Upper Stratosphere

# The QD Change in the Sum of all $O_3$ Losses (NO<sub>x</sub>, ClO<sub>x</sub>, HO<sub>x</sub>, etc.)

is driven by the NO<sub>x</sub> Loss Changes in Baseline

Fixed Dyn: QD  $O_3$  loss changes are dominated by Cl changes



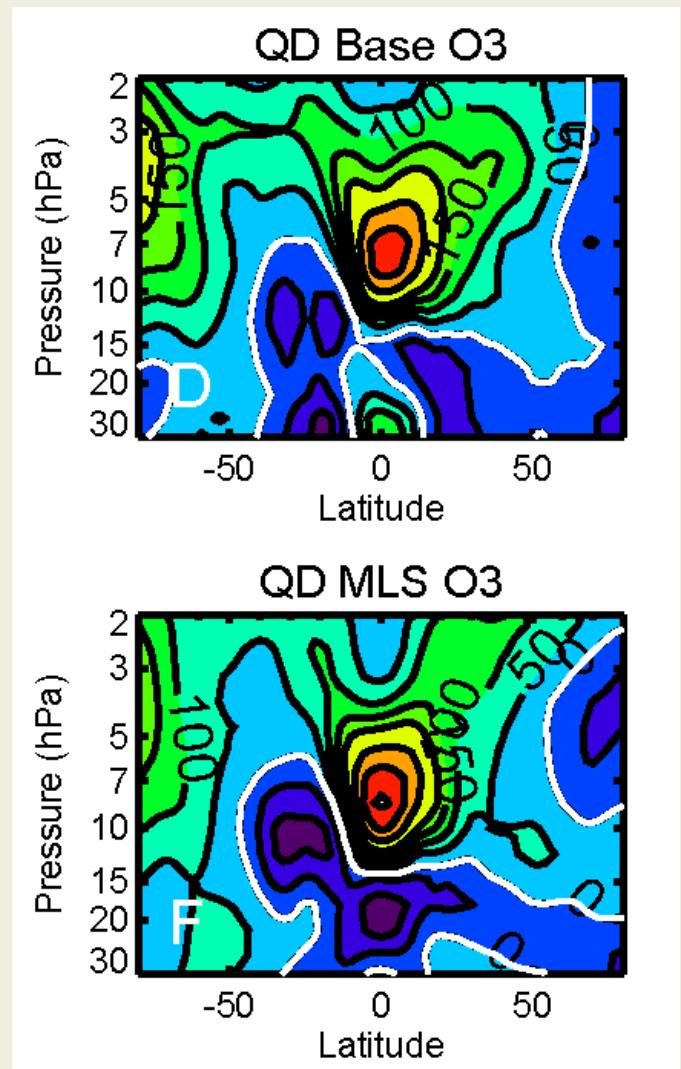
BASELINE:  
QD  $O_3$  loss changes are dominated by NO<sub>x</sub> changes. ClO<sub>x</sub> changes contribute ~3 hPa and above.

And the denouement....BASELINE QD O<sub>3</sub> changes (ppb) look very much like observed O<sub>3</sub> changes 2005-2021

Arctic Upper Strat: O<sub>3</sub> decreases! It's not a lot (<200 ppb) but it's not increasing. Driven by NO<sub>x</sub>

The Tropics: middle strat has a big O<sub>3</sub> increase because increased upwelling reduced NO<sub>x</sub> and the loss it causes

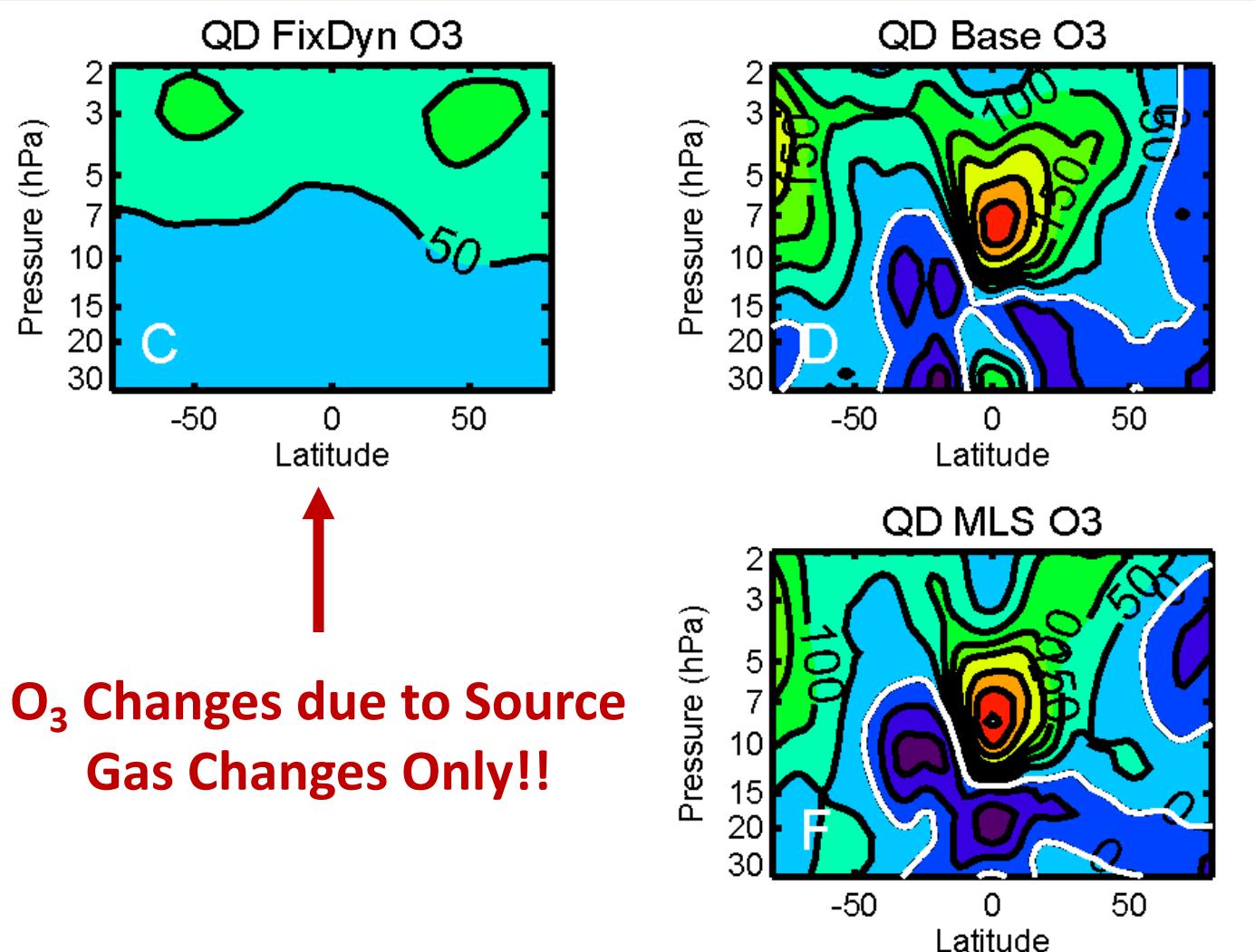
Antarctic Upper Strat: O<sub>3</sub> increases – reduced loss by ClO<sub>x</sub> and less loss by NO<sub>x</sub> (15-5 hPa)



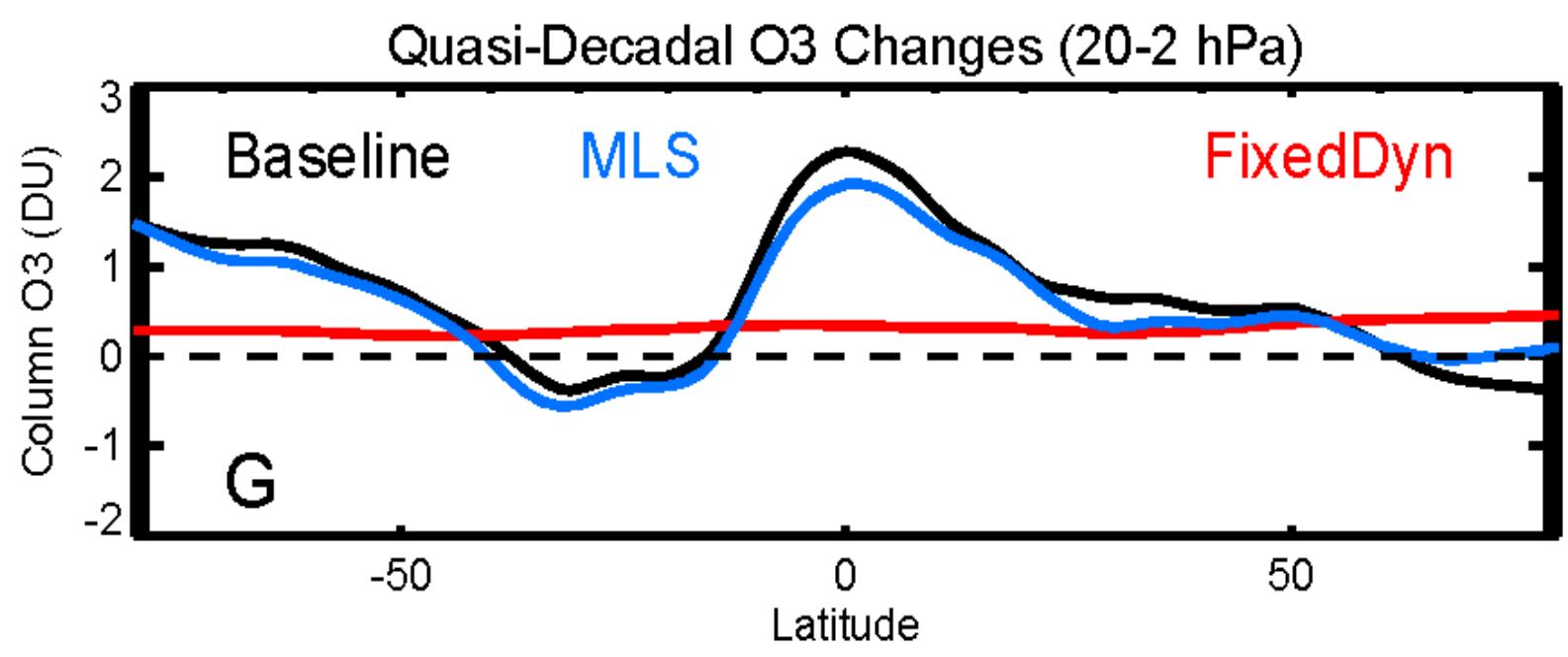
And the denouement....Fixed Dyn QD  $O_3$  changes look nothing like observed  $O_3$  changes 2005-2021

The Dynamically driven chemical changes are BASELINE – FixDyn  $O_3$  (lower left panel).

Much of the observed  $O_3$  changes 2005-2021 are caused by dynamically driven composition change!!



# Dynamically driven chemical changes from 20-2 hPa column Affect Total Column O<sub>3</sub> Trends



*Check out the excellent agreement between MLS (blue) and the GMI CTM (black)*

The 20-2 hPa column O<sub>3</sub> Quasi-Decadal changes caused by dynamics:

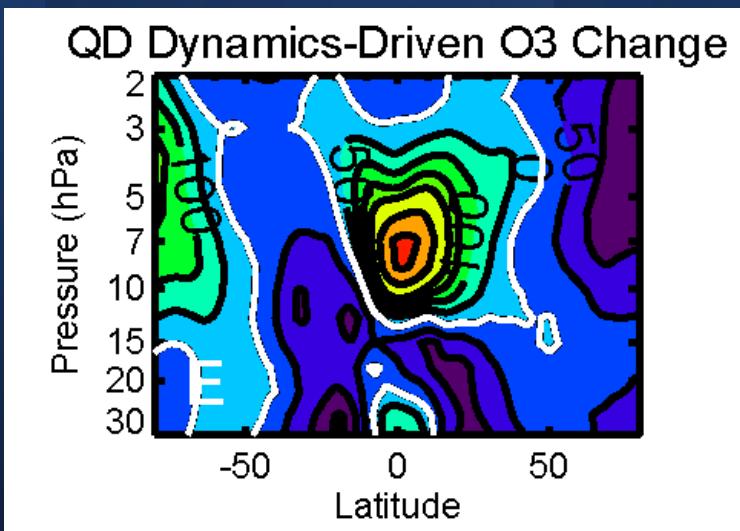
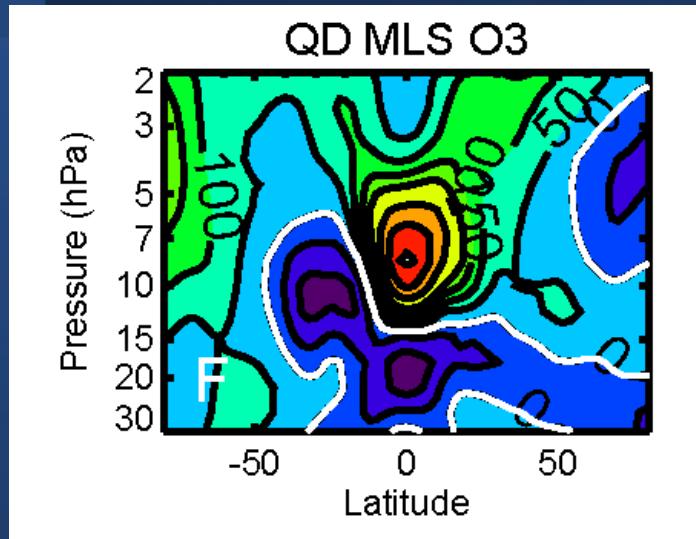
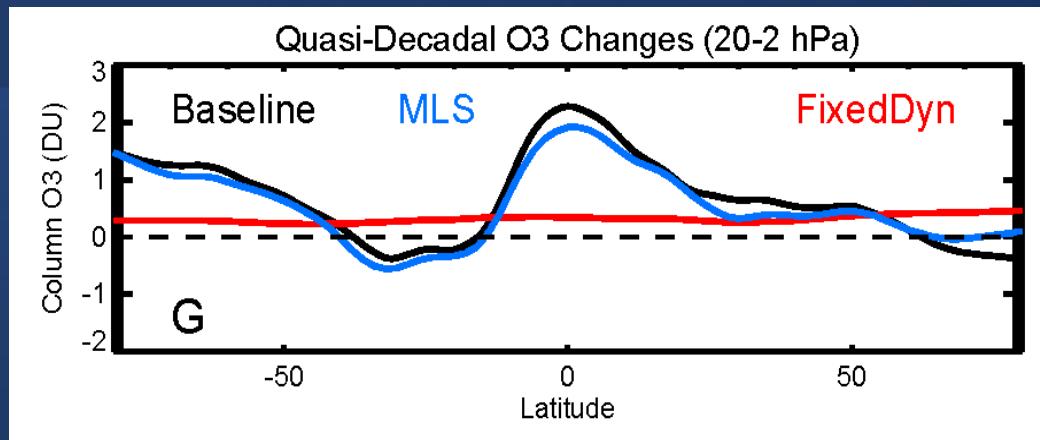
+2 DU in the tropics, -0.5 DU north of 60°N, and +1 DU in the Antarctic.

## Questions/Thoughts...

Will QBO changes persist? Is this a trend or just variability?

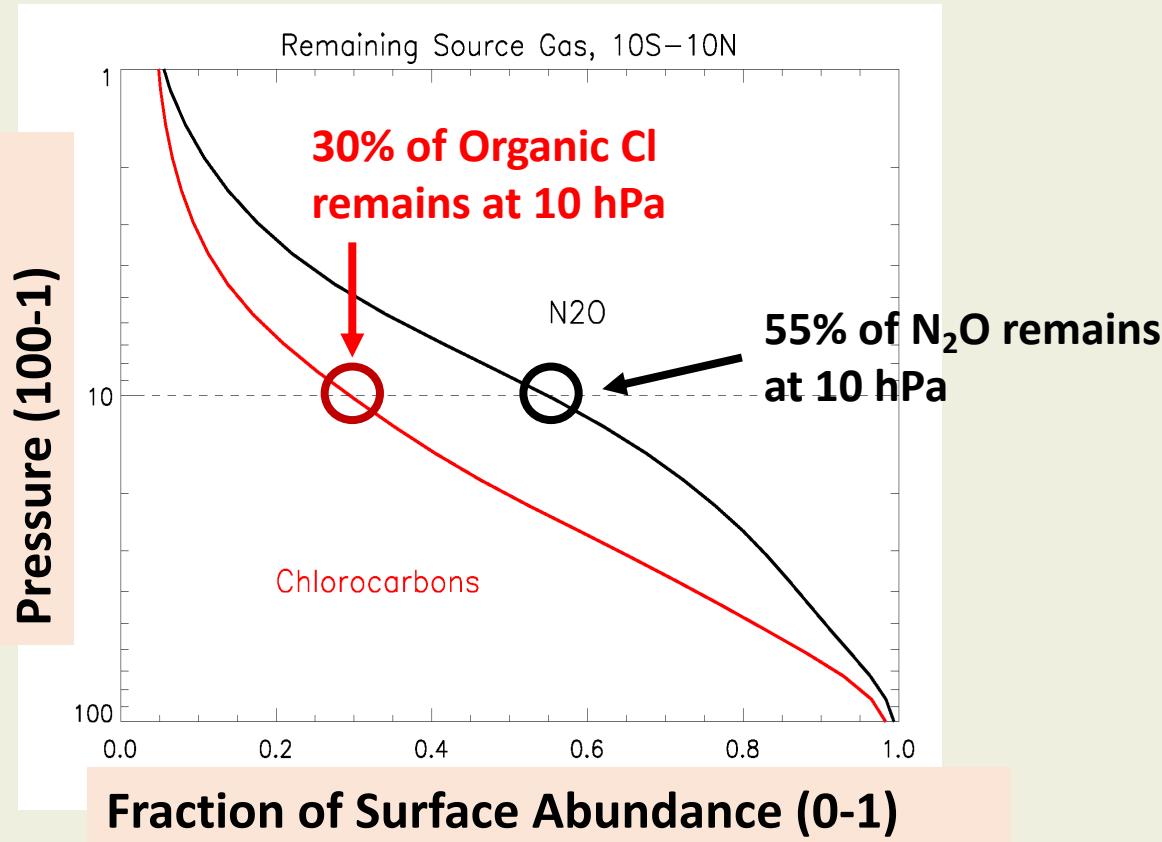
Chemistry Climate Models parameterize the QBO and can't respond physically to changes in forcing. They can't produce/predict this kind of effect on O<sub>3</sub>.

O<sub>3</sub> trend regressions fit the QBO with terms for the QBO 30 and 50 hPa (transport!). This can't regress changes in composition that affect O<sub>3</sub> chemistry



Please check out our 2022 GRL paper (Strahan et al.)

# Why is there hemispheric asymmetry in Loss by $\text{NO}_x$ but not the loss by $\text{ClO}_x$ ?

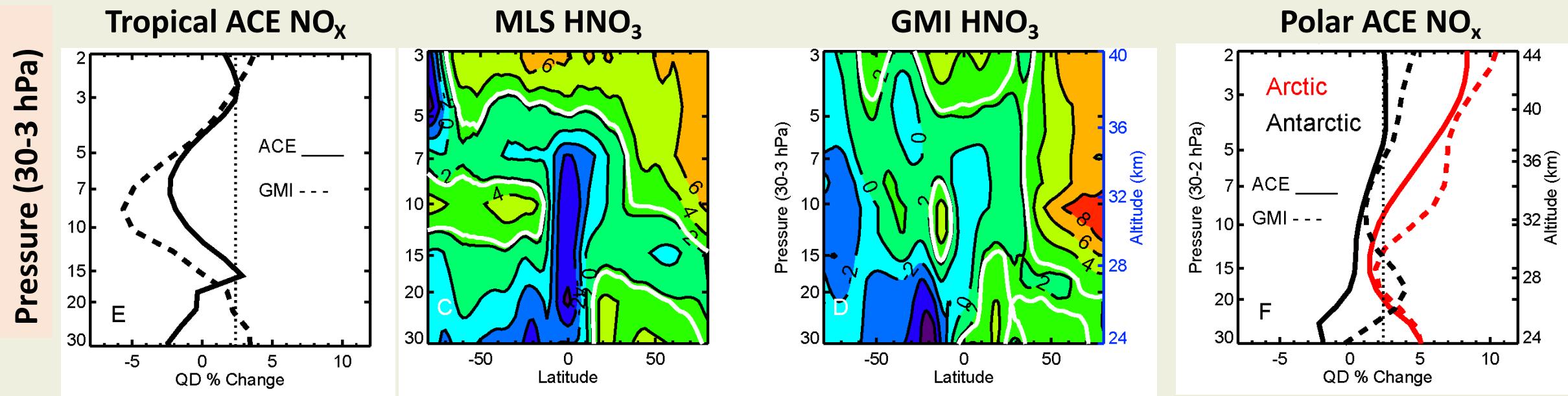


**Upwelling changes affect product gases (e.g., radicals) by changing source gas distributions**

**The upwelling changes are 10 hPa and above**

**The CFCs are mostly photolyzed below 10 hPa but most N<sub>2</sub>O is still unreacted.**

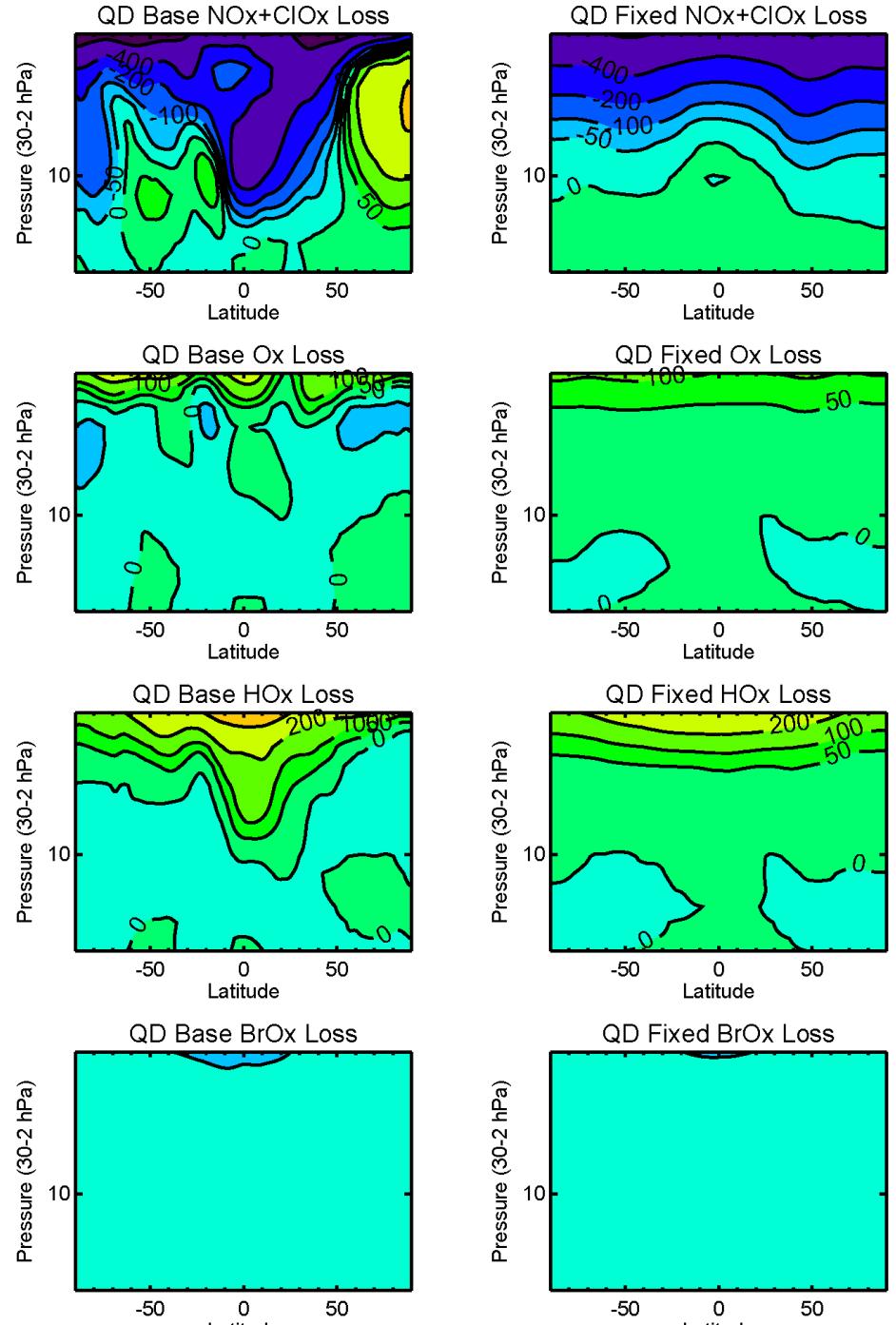
# Chemical impacts of Circulation Change: Increased Net Production of $\text{NO}_y$ is transported to the Arctic



In the tropics,  $\text{N}_2\text{O}$  is relatively long-lived below  $\sim 7 \text{ hPa}$ , thus  $\text{N}_2\text{O}$  and  $\text{NO}_y$  are largely transport controlled.  
 $\text{N}_2\text{O}$  and  $\text{NO}_y$  are anti-correlated

# All O<sub>3</sub> Loss Cycles (QD Change), Fixed Dyn (left) and Baseline (right)

NO<sub>x</sub> + ClO<sub>x</sub>  
Ox  
HO<sub>x</sub>  
BrO<sub>x</sub>



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